Supporting Information for "Gravity, topography, and melt generation rates from simple 3D models of mantle convection"

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1. Figures S1 to S6

November 7, 2019, 8:48am

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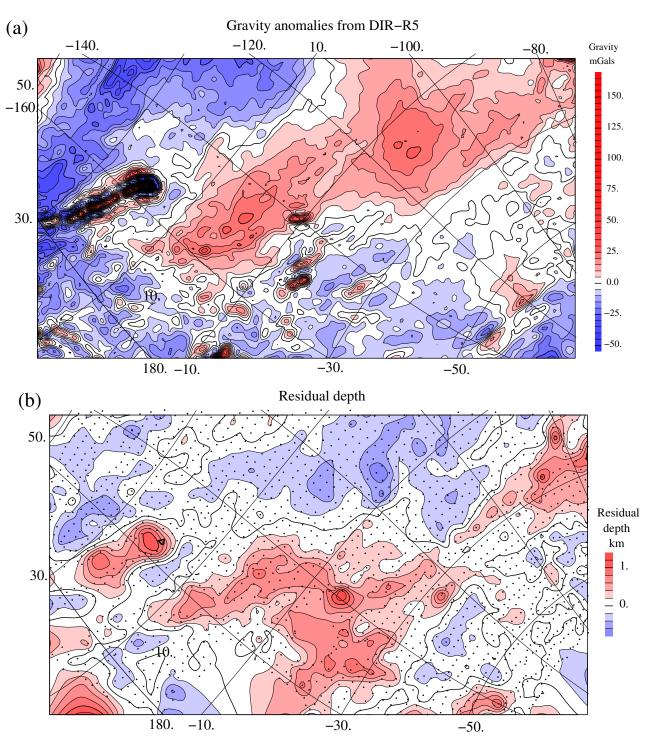


Figure S1. (a) Gravity field for the Pacific from DIR-R5, with coefficients l=2 set to 0 and a filter applied, falling to 1/2 at 250 km, to remove the short wavelength components. (b) Residual depths, averaged over $2^{\circ} \times 2^{\circ}$ boxes (Crosby et al. 2006). The dots show the locations of the resulting averages. Oblique Mercator projection with axis 40° N, -50° E.

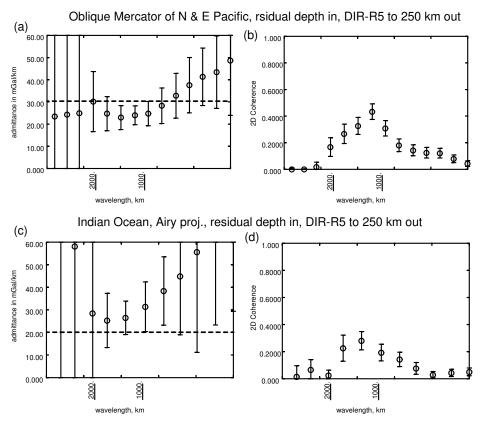


Figure S2. Admittance and coherence from the same regions used for Figure 12 (c)-(f), calculated using Hoggard et al.'s (2017) estimates of residual depth rather than those of Crosby et al. (2006).

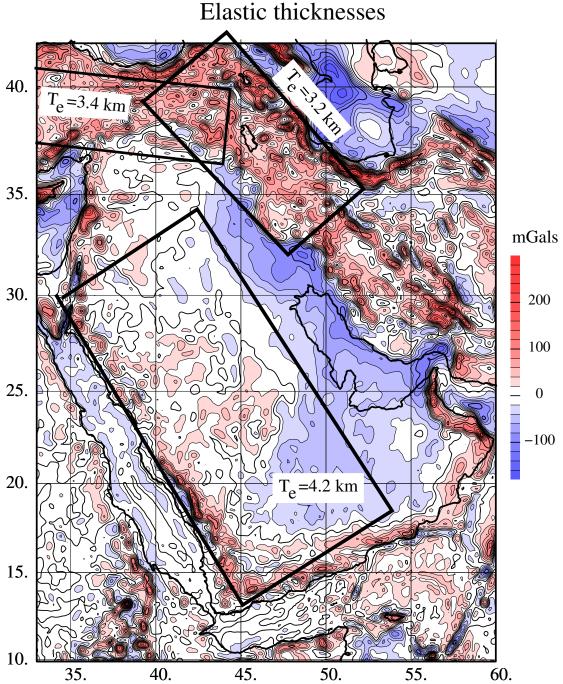


Figure S3. Boxes used to estimate the elastic thickness T_e of different parts of the Middle East, superimposed on gravity anomalies from Eigen6c (Förste et al. 2011), with the coefficients from l = 2 to 7 to 0, and applying a taper $f_{,} = (l - 7)/5$, to those from l = 8 to 11. A low pass filter falling to 1/2 at 50 km was applied to the coefficients to remove the short wavelength anomalies from uncompensated topography.

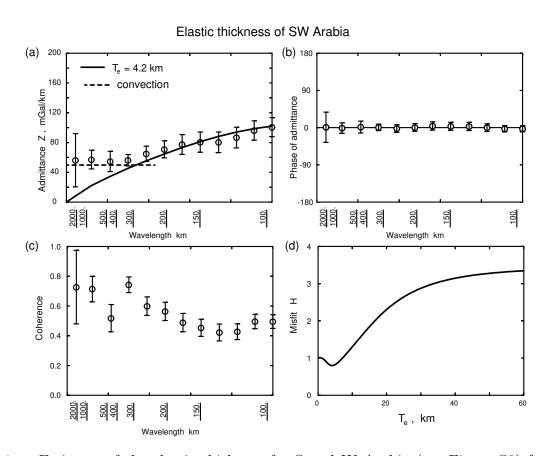


Figure S4. Estimate of the elastic thickness for S and W Arabia (see Figure S3) from the admittance, taking the topography as input, gravity from Eigen6c as output.

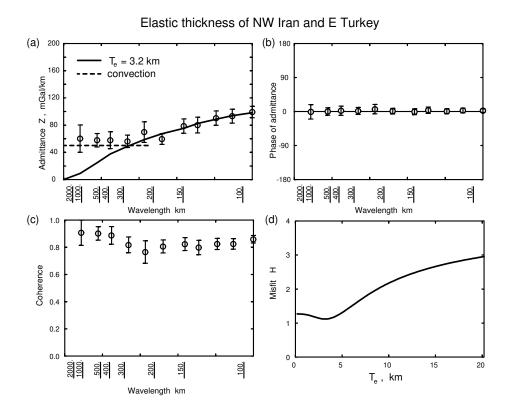


Figure S5. Estimate of the elastic thickness for E Turkey and NW Iran (see Figure S3) from the admittance, taking the topography as input, gravity from Eigen6c as output.

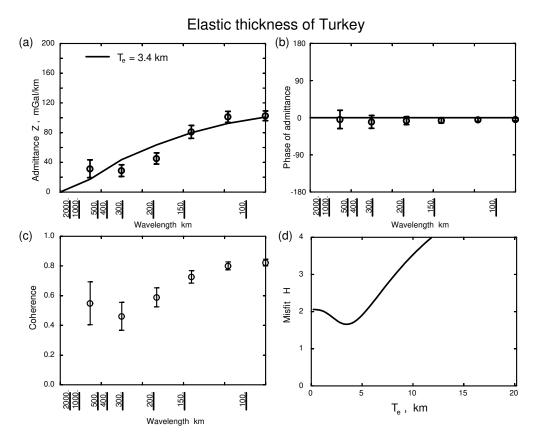


Figure S6. Estimate of the elastic thickness for Anatolia (see Figure S3) from the admittance, taking the topography as input, gravity from Eigen6c as output.